

记辽西一侏罗纪哺乳动物下颌骨

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关键词 辽宁凌源 中侏罗世 三锥齿兽目

内 容 提 要

本文记述的发现于辽宁凌源县房身的中生代哺乳动物下颌骨化石,代表了三锥齿兽目环齿兽科(?)—新的类型, *Liaotherium gracile* gen. et sp. nov.。依据无脊椎动物化石和区域地层对比,其时代似应为中侏罗世。文中还简要讨论了满洲兽和远藤兽的时代。

1976年,中国地质科学院地质研究所程政武、岳昭在辽宁西部进行中生代地层考察时,于凌源县房身附近的杂色火山碎屑岩层露头的下部,采集到一块哺乳动物的下牙床。这是辽宁省发现的第三个中生代哺乳动物化石地点。

1958年,杨钟健曾报道过凌源县鸽子洞发现的蜥蜴类的细小矢部龙 (*Yabeinosaurus tenuis* Endo and Shikama, 1942) 化石,并据此认为产此化石的地层(当时认为是“九佛堂组”)的时代为晚侏罗世(Young, 1958)。根据杨钟健文中的剖面图(P.152)和我们的两次实地核对,新发现的哺乳动物化石与矢部龙化石应该产于同一地点同一层位。

一、产化石地层的时代

化石产地房身,地理座标为北纬 $41^{\circ}0'$,东经 $119^{\circ}05'$ (图1),位于平泉盆地的南部边缘,行政上隶属于辽宁省凌源县。

化石地点附近出露的与化石层有关的地层有三套,自下而上为: 1) 前寒武纪长城系白云岩和石英岩; 2) 侏罗系紫红、棕红、灰绿色钙质、凝灰质泥岩、砂岩及砾岩夹少量安山质凝灰岩; 3) 安山质、流纹质火山角砾岩、安山岩夹岩屑、玻屑凝灰岩。长城系与上覆的侏罗系为不整合接触,侏罗纪的两套地层之间的接触关系为整合或假整合。在1:20万地质图凌源幅中(辽宁省地质局区域地质测量队,1965),该地点的这两套地层统称为蓝旗组(中侏罗统)。哺乳动物化石就产在不整合面之上的杂色火山碎屑岩的下部(图2)。

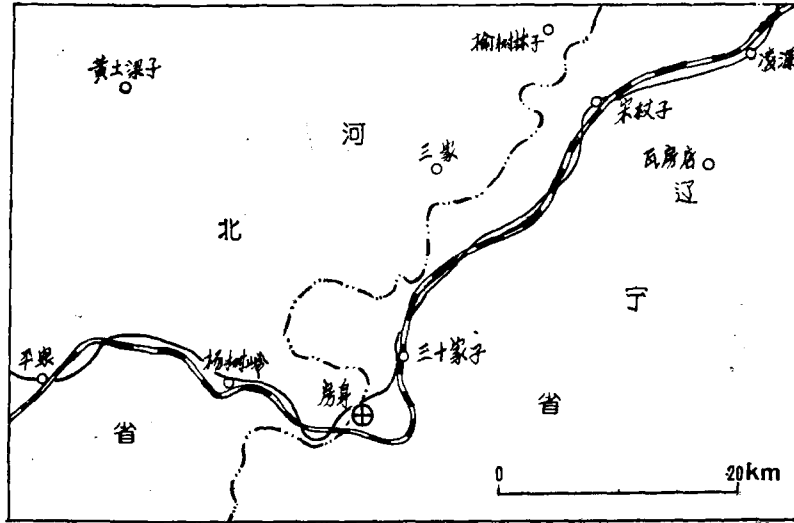


图 1 凌源县房身化石地点位置图

Fig. 1 The location of the fossil site at Fangshen, Lingyuan County, Liaoning Province

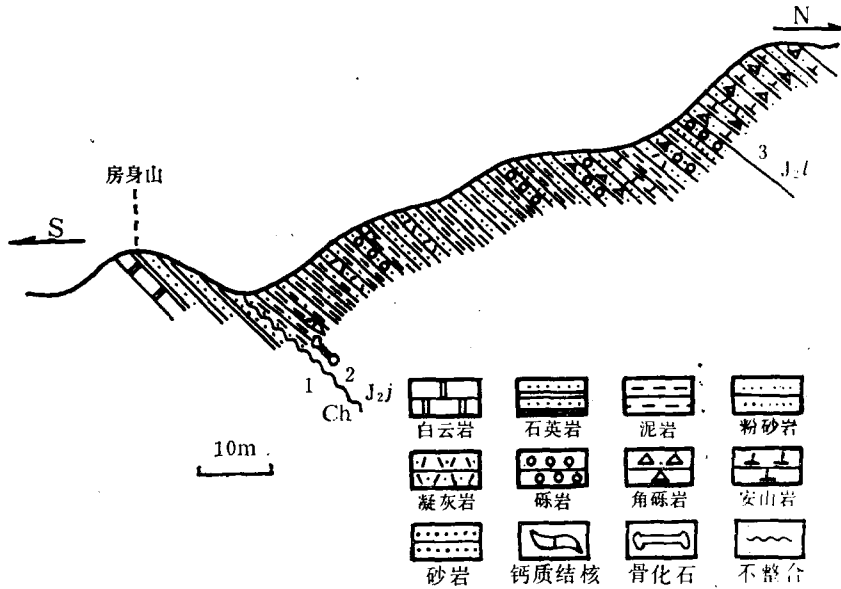


图 2 凌源房身化石地点剖面图

Fig. 2 The section of the fossil site at Fangshen, Lingyuan County, Liaoning Province

根据我们野外观察，房身一带在区域构造上属于平泉盆地。产哺乳动物化石的杂色火山碎屑岩在地层层序及岩性特征方面，均与河北平泉盆地的九龙山组一致，而与东邻的朝阳盆地中相当层位的海房沟组黄绿色地层差别明显。实际上，这套杂色火山碎屑岩地层是河北平泉地区九龙山组向辽宁西部延伸的一部分。这一点可以从 1:20 万地质图平

泉幅(河北省地质局第二区域地质测量大队, 1976)和凌源幅(辽宁省地质局区域地质测量队, 1965)中清楚地看出来。因此, 本文将凌源县房身一带的这套杂色火山碎屑岩从“蓝旗组”中划出, 并采用九龙山组一名。

在整个冀北地区(包括北京西山), 九龙山组伏于中侏罗统髫髻山组火山岩之下。辽西朝阳盆地相当层位的海房沟组伏于中侏罗统蓝旗组之下。王思恩(1984)将九龙山组中的叶肢介化石称为真叶肢介—三饰叶肢介群 (*Euestheria-Triglypta* Fauna), 认为其时代是中侏罗世早期。洪友崇(1984)依据九龙山组的昆虫化石建立了九龙山组昆虫组合 (*Jiulongshan Insecta Assemblage*), 隶属于燕辽昆虫群 (*Yenliao Insecta Fauna*), 时代定为中侏罗世。王东方等(1984)对辽西侏罗—白垩系的火山岩进行了系统的同位素年龄测定, 测得蓝旗组样品的 K-Ar 同位素年龄为 158.1 ± 8 百万年, 认为其时代可能是中侏罗世晚期到晚侏罗世早期。因此, 位于其下的海房沟组(与九龙山组层位相当)的时代应属中侏罗世。

以上古生物和同位素年龄测定的资料表明, 凌源县房身含哺乳动物化石地层的时代应属中侏罗世。这样, 如果杨钟健 (Young, 1958) 研究报道的与房身哺乳动物化石发现于同一地点同一层位的蜥蜴类化石确实是矢部龙 (*Yabeinosaurus tenuis* Endo and Shikama) 的话, 那么义县枣茨山和凌源鸽子洞(即房身)两地点的矢部龙化石就应属于两个不同的层位: 前者是上侏罗统义县组, 后者为中侏罗统九龙山组。

二、化石记述

哺乳动物纲 *Mammalia* Linnaeus, 1758

原兽亚纲 *Prototheria* Gill, 1872

始兽次纲 *Eotheria* Kermack and Mussett, 1958

三锥齿兽目 *Triconodonta* Osborn, 1888

环齿兽科(?) ? *Amphilestidae* Osborn, 1888

辽兽(新属) *Liaotherium* gen. nov.

属型种 纤细辽兽(新属新种) *Liaotherium gracile* gen. et sp. nov.

特征 同属型种。

名称来源 “辽”(Liao) 是化石产地辽宁省的简称。

纤细辽兽(新属新种) *Liaotherium gracile* gen. et sp. nov.

(图 3)

正型标本 较完整的右下牙床一块, 具最后一枚臼齿($M_5?$)。中国地质博物馆标本编号: V2006。

产地及层位 辽宁凌源县房身, 中侏罗统九龙山组。

特征 下颌骨水平枝长而纤弱; 骨体上、下缘平直, 近于平行; 下缘在下颌骨前端呈弧形上翘。冠状突小而弱; 关节突略高于齿列; 髁上凹发育。齿式可能为 $I_{2+1}C_1P_3M_5$ 或 $I_{2+1}C_1P_4M_0$ 。

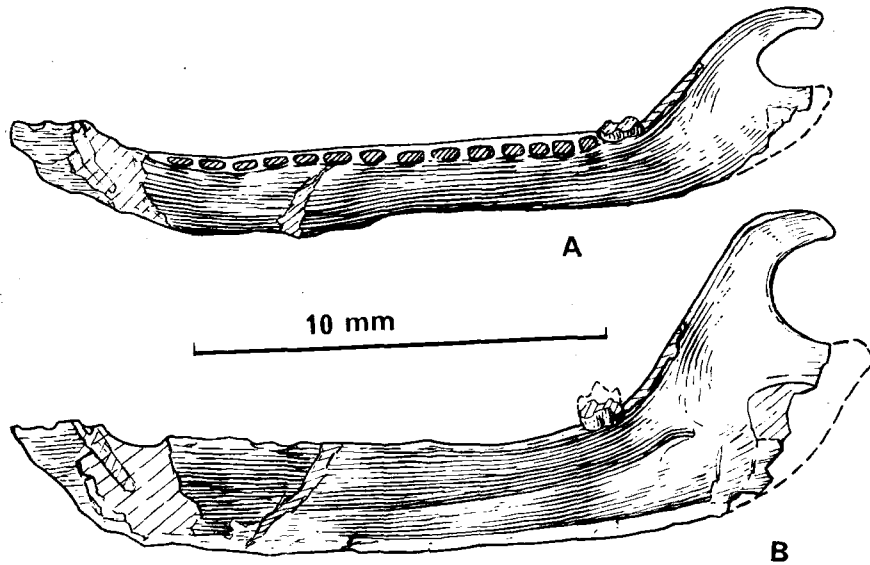


图3 纤细辽兽(新属新种)的右下颌骨 (V2006)

Fig. 3 The right lower jaw of *Liaotherium gracile* gen. et sp. nov. (V2006)

A. 内侧顶视 (internal-top view); B. 内侧视 (internal view)

名称来源 Gracil-(拉), 纤弱的、纤细的, 表示该下颌骨比较纤弱。

描述 下颌骨除关节突部分破损外, 余皆保存完整, 长 22.2 毫米, 高 8.6 毫米。水平枝较长而相对纤弱, 长度大于 15.6 毫米。下颌骨体高度前部 2.7 毫米, 后部 2.4 毫米, 上、下缘平直, 近于平行, 下缘前端向上呈弧形弯曲, 上缘厚度约 1 毫米。除最后一枚臼齿外, 其余牙齿均未保存。在最后一枚臼齿之前, 有 14 个颊齿窝, 齿窝近方形。齿窝大小、间距无明显的变化。从第一个颊齿窝前缘到最后一枚臼齿后缘长约为 12.6 毫米。最后一枚臼齿靠近上升枝, 长约 1.1 毫米。下颌骨前端还保存有两个门齿窝, 在它们与颊齿窝之间, 有一段骨体破损, 可见一较大的齿窝痕迹, 可能是犬齿窝所在的位置。下颌骨上升枝高约 5.8 毫米, 与水平枝之间的夹角约为 120° 。冠状突小而弱, 关节突破损, 从留下的印痕上看, 它略高于下齿列。髁上凹发育。

讨论 由于保存的唯一一枚牙齿在采集时被损坏, 因此很难对这一化石进行更详细的描述并确定其系统分类位置。但根据在野外的初步观察, 最后一枚臼齿 ($M_3?$) 具有前后排列的三个齿尖, 中间的主尖比前后两个齿尖大而且高。这一特点表明它与产于英国 Stonefield 中侏罗统中的三锥齿兽类的环齿兽 (*Amphilestes*) 比较接近。如果它的颊齿象多数三锥齿兽类一样, 均具双齿根的话, 这一化石的齿式就可能为 $I_{2+7}C_1P_3M_5$ 或 $I_{2+7}C_1P_4M_4$ 。

虽然 *Triconodon* 的颊齿齿式为 P_4M_4 , 但是 *Triconodon* 以及其它 Triconodontidae 科成员的颊齿明显地从前往后逐渐变大, 且臼齿的三个齿尖近于等大, 与凌源标本明显不同。凌源标本的颊齿形态与环齿兽科 (*Amphilestidae*) 成员相近, 基于这一点, 本文暂将它归入 *Amphilestidae* 中。

环齿兽科是 Osborn 1888 年命名的, 置于三锥齿兽目 (*Triconodonta*) 中, 当时仅

作为一个亚科。Kühne 1958 年将它提升为科。Jenkins and Crompton (1979) 也认为环齿兽科应放在三锥齿兽目中。但 Mills (1971)、Crompton (1974) 和 Freeman (1979) 对此提出不同看法, 认为环齿兽科应归入对齿兽目 (Symmetrodonta Simpson, 1929)。我们认为, 环齿兽类颊齿的三个齿尖前后排列, 与对齿兽类齿尖排列成三角形明显不同, 而与其他三锥齿兽类的齿尖排列方式一致。因此, 将其归入三锥齿兽目中可能更为合适。

环齿兽科已描述的共有四个属, 它们是英国中侏罗世 *Amphilestes* Owen, 1871 和 *Phascolotherium* Owen, 1838, 美国怀俄明州晚侏罗世 *Phascolodon* Simpson, 1925 和 *Aploconodon* Simpson, 1925。 *Phascolodon* 和 *Aploconodon* 仅保存了下颌骨后部, 无法与凌源的标本详细比较, 但它们的最后一枚臼齿明显比前面的臼齿小, 仅 0.7 毫米长, 而凌源标本上最后一枚臼齿长约 1.1 毫米。 *Phascolotherium* 的下颌骨体上缘较平直, 但其下缘呈弧形, 冠状突粗壮, 且齿式为 $I_4C_1P_2M_5$, 这些均与凌源的标本不同。 *Amphilestes* 与凌源的标本在以下几个方面比较接近: 下颌骨体上缘较平直, 关节突略高于齿列。但 *Amphilestes* 下颌骨较粗壮、骨体下缘呈弧形、冠状突大而粗壮, 颊齿齿式为 P_4M_5 , 明显地区别于凌源的标本。

克拉美丽兽 (*Klamelia zhaopengi* Chow and Rich, 1984) 发现于我国新疆准噶尔盆地北部克拉美丽地区中侏罗世晚期或晚侏罗世早期地层中。这是我国第一次发现的三锥齿兽类化石。从保存的部分看, 凌源标本与它之间差别明显。Chow and Rich (1984) 将 *Klamelia* 与蒙古早白垩世 *Gobiconodon* Trofimov, 1987 和 *Guchinodon* Trofimov, 1978 合在一起, 建立了一个新亚科——戈壁兽亚科 (*Gobiconodontinae*), 置于环齿兽科中。Jenkins and Schaff (1988) 在研究美国蒙大拿州早白垩世 *Gobiconodon ostromi* (=undescribed amphilestid of Jenkins and Crompton, 1979, from Cloverly Formation) 时指出, *Gobiconodon* 门齿增大、犬齿小, 与一般的环齿兽类明显不同, 因而建立一新科——戈壁兽科 (*Gobiconodontidae*), 并将 *Klamelia* 排斥在新科之外。他们还认为 *Klamelia* 与其它环齿兽类也有明显的差别, 可能不具有很近的亲缘关系。从 Trofimov (1981) 和 Jenkins and Schaff (1988) 文章的插图、图版及描述上看, 凌源的标本与 *Gobiconodon* 和 *Guchinodon* 之间存在明显的差异, 主要表现在后两属个体较大、下颌骨粗壮、骨体上、下缘均呈弧形, 臼齿前牙齿数目少, 齿式为 $I_1C_1P_{3-4}M_5^{1)}$, 前臼齿均为单齿根等方面。另外, *Gobiconodon* 的关节突略低于下齿列, 冠状突较强壮, 也与凌源标本不同。

1986 年, Bonaparte 根据阿根廷 Patagonia 晚白垩世的一枚左 M_1 建立了 *Austrotriconodon mckennai*, 归入三锥齿兽科中 (Bonaparte, 1987)。因它仅有一枚左 M_1 , 凌源标本无法与之对比, 但其 M_1 长达 2.8 毫米, 可见其个体比 V2006 大得多。

根据以上的讨论, 我们认为凌源的标本似应代表了环齿兽科一个新属种, 命名为纤细辽兽 (*Liaotherium gracile* gen. et sp. nov.)。

产于辽宁东部与凌源标本层位相近的满洲兽 (*Manchurodon simplicidens* Yabe and Shikama, 1938), 个体较大, 下颌骨粗壮, 水平枝往前逐渐变细, 骨体上、下缘不平行, 齿

1) Trofimov (1981) 认为其齿式为 $I_3C_1P_4M_5$, 但插图中在臼齿前仅有 6 个齿窝。

表 1 辽宁侏罗纪—早白垩世含脊椎动物化石地层对比表
Table 1 The correlation of vertebrate-bearing Jurassic and Lower Cretaceous in Liaoning

		冀北—辽西地区		辽东地区	
		平泉—凌源盆地 ¹⁾		朝阳盆地 ²⁾	
		孙家湾组	阜新组	九佛寺组	义县组
白垩系 K	下统 K ₁			<i>Endotherium Teilhardosaurus</i>	
		九佛寺组 (南店组)	<i>Lycoptera</i>	<i>Pittacosaurus Lycoptera Peipiaosteus</i>	普兰店组
	上统 J ₃	义县组 (花营组、西瓜园组)	<i>Lycoptera Rhynchosaurus³⁾</i>	<i>Lycoptera, Peipiaosteus, Mancharochelys, Yabeinosaurus</i>	桂云花组
侏罗系 J		土城子组 (后城组)	<i>Xuanhuasaurus⁴⁾</i>	<i>Chaoyuungosaurus Palaeoniscoid Jeholosauripus</i>	
	中统 J ₂	蓝旗组 (髫髻山组)			
		九龙山组	<i>Liaotherium gen. nov., Yabeinosaurus Sauropoda</i>	<i>Palaeoniscoid</i>	瓦房店组
	下统 J ₁				
下伏地层		长城系	北票组 兴隆沟组 坤头波罗组		寒武系

1) 除“九龙山组”引用河北省的组名外,其余均沿用辽西的组名。

2) 包括阜新地区。

3) 经考察,该属产于凌源大南沟,所谓“九佛寺组”实为义县组(包括金剛山组和吐噜组)。

4) 该属产于河北宣化堰家沟。

式为 $I_1C_7P_4M_4$, 具明显的齿隙, 臼齿除主尖外, 仅主尖之后有一小齿尖, 被归入对齿兽目 *Amphidontidae* 科中。凌源的标本与之相比差别很大。

三、满洲兽与远藤兽的时代

满洲兽和远藤兽是发现于我国东北地区的两种中生代哺乳动物, 关于它们的时代一直存在争议。

满洲兽是矢部长克和鹿间时夫根据发现于辽宁复县(当时称瓦房店)砬子窑煤矿的一个下牙床命名的, 原作者认为其时代为中侏罗世 (Yabe and Shikama, 1938)。此后, 关于它的时代长期存在争议。Teilhard de Chardin and Leroy (1942)、周明镇 (1953) 和 Patterson (1956) 都认为其时代为早白垩世。张法奎 (1984) 认为其时代为晚侏罗世。

根据区调成果, 瓦房店、普兰店一带的中生代地层层序已经建立 (表 1, 复县盆地)。*Manchurodon* 产于瓦房店组含煤地层中(过去曾误认为属阜新煤系)。瓦房店组中所产的双壳类和植物化石多为早、中侏罗世的分子。从岩相古地理和古构造运动方面考虑, 我们认为瓦房店组大体与海房沟组相当, 其时代应属中侏罗世。其中所含的 *Manchurodon* 亦应为中侏罗世哺乳动物的化石代表之一。

远藤兽 (*Endotherium niinomii* Shikama, 1947) 是鹿间时夫命名的, 标本产地是阜新新邱煤矿, 原作者认为其时代为晚侏罗世 (Shikama, 1947)。1953 年, 周明镇指出 *Endotherium* 与重褶齿獾 *Zalambdalestes* 接近, 其时代可能是早白垩世。

通过实地考察, 证实远藤兽产于阜新组煤系地层中。过去对该组的时代存在较大的争议。近年来, 虽然对侏罗—白垩系界线划分意见仍有分歧, 但随着新材料的积累和研究的深入, 关于阜新组的时代却取得了比较一致的意见。依据对阜新组植物化石 (张志诚, 1987) 及下伏沙海组叶肢介化石 (王五力, 1987) 的研究, 上述两组的时代均属早白垩世。因而, 可以认为产于阜新组中的远藤兽化石的时代为早白垩世。

到目前为止, 辽宁省侏罗纪—早白垩世地层中已发现了不少脊椎动物化石, 包括鱼类、爬行类及哺乳类, 含脊椎动物化石层位及地层对比关系见表 1。

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A MAMMALIAN LOWER JAW FROM THE JURASSIC OF LINGYUAN, LIAONING

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Key words Lingyuan, Liaoning; Middle Jurassic; Triconodonta

Summary

In 1976, a right mandible of a Mesozoic mammal was collected by Messrs. Cheng Zhengwu and Yue Zhao of Institute of Geology, the Chinese Academy of Geological Sciences, from the Middle Jurassic in Lingyuan County, Liaoning Province, North-east China (fig.1), when they were studying the Mesozoic stratigraphy of that area. This fossil site is most probably from the same locality as that of *Yabeinosaurus tenuis* reported by C.C. Young in 1958. Both fossils are derived from the same horizon.

The mammalian fossil here described occurs in the Jiulongshan Formation, a set of sediments of purplish red, brownish and greyish green calcareous and tuffaceous mudstone, sandstone and conglomerate, with intercalating andesitic tuffaceous layers, superposing on beds of dolomite and quartzite of Changcheng System (Pre-Cambrian) and overlain by the andesitic and rhyolitic volcanic breccia and andesite of Lanqi Formation (fig.2).

According to the studies of fossil conchostraca (Wang, S.E., 1984) and insects (Hong, 1984) collected from Jiulongshan Formation, its geological age is tentatively considered to be Middle Jurassic. While, the K—Ar radiometric age for the overlying Lanqi Formation is about 158.1Ma (Wang, D.F., 1984).

Systematic Paleontology

Class Mammalia Linnaeus, 1758

Subclass Prototheria Gill, 1872

Infraclass Eotheria Kermack and Mussett, 1958

Order Triconodonta Osborn, 1888

Family ?Amphilestidae Osborn, 1888

Genus *Liaotherium* gen. nov.

Type species *Liaotherium gracile* gen. et sp. nov.

Diagnosis As for the type species.

Etymology "Liao" is the abbreviation for Liaoning Province where the fossil was collected.

***Liaotherium gracile* gen. et sp. nov.**

Holotype The Geological Museum of China, Beijing, specimen no. V2006, a right mandible with the last molar ($M_5?$) (fig.3).

Locality and Stratigraphical Horizon Fangshen, Lingyuan County, Liaoning Province (approximately 41°N , $119^\circ05'\text{E}$), from the lower part of Jiulongshan Formation, Middle Jurassic.

Diagnosis The horizontal ramus of the mandible long and slender; both the upper and lower borders of the body straight and nearly parallel; the lower border curves slightly upwards at the anterior end. Coronal process small and feeble; articular process a little above the top of the lower dentition; supracondyloid notch well-developed. Lower dentition is inferred to be: $I_{2+7}C_1P_3M_5$ or $I_{2+7}C_1P_4M_4$.

Etymology *Gracil-* (Latin), thin, in reference to the slenderness of the mandible.

Description A complete mandible, broken at the articular process, about 22.2 mm long and 8.6mm high. The horizontal ramus is long, more than 15.6mm, and quite slender. The height of the body of mandible is 2.4—2.7mm. It is a little higher at the anterior end than at the posterior one. The lower and upper borders are straight and nearly parallel. The lower border curves upwards at the anterior end of the mandible. The upper border is about 1mm in width. Except the last molar, all other teeth are absent. There are 14 cheek tooth alveoli before the last molar. These alveoli are approximately square. The length of the cheek tooth dentition is about 12.6 mm. The last molar is close to the uplifting branch and about 1.1mm in length. At the anterior end of the mandible, two incisor alveoli are preserved. The damaged part between the incisor and the cheek tooth alveoli may be the position of the canine. The uplifting branch of the mandible is 5.8mm high and at an angle of 120° to the horizontal ramus. The coronal process is small and feeble. Supracondyloid notch is well-developed. The articular process is broken. Its print on the rock shows that it is a little higher than the lower dentition.

Remarks As the last molar, the only preserved tooth, was damaged, it is difficult to ascribe it to an appropriate systematic position. According to the field record, as was observed in the field when it was first exposed, this molar has three cusps arranging in a line. The main middle cusp (a) is larger and higher than the other two cusps (b and c). This and its other dental features show that the fossil mammal from Lingyuan, Liaoning is quite possibly similar to *Amphilestes* Owen, 1871 from Stonefield Slate of Middle Jurassic, Oxfordshire, England, so that we tentatively referred this specimen to the family Amphilestidae Osborn, 1888. If all its premolars are double-rooted, its lower dentition is possibly $I_{2+7}C_1P_3M_5$ or $I_{2+7}C_1P_4M_4$.

Known genera of Amphilestidae now include: *Amphilestes* Owen, 1871 and *Phas-*

colotherium Owen, 1838 from Middle Jurassic of England, *Phascolodon* Simpson, 1925 and *Aploconodon* Simpson, 1925 from Late Jurassic of Wyoming, USA.

Aploconodon and *Phascolodon* could not be compared in detail with the specimen described in this paper because only fragments of posterior parts of their mandibles were known. Their last molars, 0.7mm in length, are obviously smaller than the molars anterior to them, while the last molar of Lingyuan specimen is 1.1mm long. *Phascolotherium* differs from V2006 in the curved lower border of mandible, sturdy coronal process, and dentition ($I_4C_1P_2M_5$). *Amphilestes* was similar to V2006 in having straight upper border of mandible, and articular process somewhat higher than the lower dentition, but it obviously differed from the latter in having thick-set mandible with a curved lower border, sturdy coronal process, and the different cheek tooth dentition (P_4M_5), etc.

Another triconodont, *Klamelia zhaopengi* Chow and Rich (1984) from late Middle or early Late Jurassic of Zungaria Basin, from which V2006 distinctly differed, was first referred to Amphilestidae. Jenkins and Schaff (1988) suggested that it does not show close relationship to the other amphilestids.

As above, the features of V2006 show that it very likely represents a new form of Amphilestidae.

On the age of *Manchurodon simplicidens* and *Endotherium niinomii*

Manchurodon simplicidens was named and described by Yabe and Shikama based on a fragmentary mandible found at Zhaziyao (Sakusiyō) Coal Mine of Fuxian in 1938. They deemed its geological age to be Middle Jurassic. Later, the age was thought to be as Early Cretaceous by Teilhard de Chardin and Leroy (1942), Chow (1953) and Patterson (1956), and Late Jurassic by Zhang, F.K. (1984).

The specimen of *M. simplicidens* was collected from coal-bearing Wafangdian Formation which had been previously mistaken for "Husin Series" (Yabe and Shikama, 1938). Most of the plant and bivalve fossils collected from this formation are the elements common in Early and Middle Jurassic. Based on stratigraphical relationship, the age of Wafangdian Formation is more probably to be Middle Jurassic, and *M. simplicidens* is a representative of a Middle Jurassic mammal.

Endotherium niinomii, described by Shikama in 1947, was from Fuxin (Husin) Formation at Xinqiu (Hsinchiu) Coal Mine in Fuxin. The original author considered that its geological age was Late Jurassic.

Chow (1953) suggested that *Endotherium* was quite advanced and seems to be similar to *Zalambdalestes*, so that it may be Early Cretaceous in age. The recent studies of fossil plants by Zhang, Z.C. (1987) from Fuxin Formation and conchostraca by Wang, W.L. (1987) from the underlying Shaihai Formation all show that Fuxin Formation should be the deposits of Early Cretaceous.

The vertebrate-bearing Jurassic-Early Cretaceous stratigraphic sequences and correlation in Liaoning are summarized in Table 1.